

# Overview of the AIRS Science Team Algorithm.



by

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**Mar 21, 2005**

**Complete Notes on remote sounding and the Atmospheric InfraRed Sounder (AIRS) Science Team retrieval methodology are on:**

**ftp ftp.orbit.nesdis.noaa.gov**

**cd pub/smcd/spb/cbarnet/reference**

rs\_notes.pdf      remote sounding notes & retrieval theory

phys640\_s04.pdf    computational methods with chapters on  
linear and non-linear least squares

airsb\_code.pdf     documentation on the code itself

# Organizations That Have Contributed to AIRS



**AIRS-AMSU-HSB**



**Atmospheric Infrared Sounder**



**Advanced Microwave Sounding Unit**



**Humidity Sounder for Brazil**



## Overview of This Talk

- Introduction to regularization
  - How the theory of remote sensing has evolved.
  - What is the basis for the AIRS methodology.
    - \* Cloud clearing philosophy in an integral component
    - \* Minimization approach is optimized for hyper-spectral sounders.
- Details of the AIRS Science Team implementation.
  - Justification for use of channel sub-sets.
  - Justification for use of trapezoids.
  - Justification for use of finite differences.
- Some post-launch issues
  - Microwave side-lobe corrections
  - Tuning versus Error Term Experiments
- A few details on the code & development system

## Need to Use “Snake-eyed Concentration



In the 1970 movie, Little Big Man, the character Jack Crabb learns from his sister than “it is possible to shoot a gun before you touch it; however, it takes lightnin’ reflexes and considerable snake-eyed concentration.”



In this context:

It is possible to understand the methodology without understanding the details of the equations; however, it still takes lightnin’ reflexes and considerable snake-eyed concentration.

## Notation, a.k.a., How I Learned to Embrace my OCD

- I adopt a notation of linear algebra that denotes the dimensions and row/column indices of the matrices. Yes, there is medication for this.
- For example, the kernel function,  $K$ , is a rank-2 matrix,  $K(n, L)$  where  $n$  is a channel index and  $L$  is a parameter index (our pressure grid).
- I will write this matrix as  $K_{n,L}$ . Having explicit indices makes transposes,  $K_{L,n}^T$ , and inverses,  $K_{L,n}^{-1}$ , more obvious.
- It is also useful when programming. Loop indices,  $n, i, j, k, L, m$  etc. used in the FORTRAN code correspond to the matrix indices in the theory documents. Obviously, I am now off the medication!
- In this notation an implied summation rule is imposed

$$\begin{aligned} \mathbf{y}_n^T \cdot \mathbf{y}_n &= C, \quad \text{a scalar} \\ \mathbf{y}_n \cdot \mathbf{y}_n^T &= C_{n,n}, \quad \text{a rank 2 matrix} \end{aligned}$$

- I also use a superscript  $i$  to denote items that are iterated. The concept of iteration evolves during the talk, so some items will begin without the iteration superscript and then will have it later.

## Least Squares (LSQ) Solution

- If we want to solve a linear system of equations of observations  $y_n$ , where  $n$  is number of channels using  $L$  geophysical parameters,  $X_L$ , we could write the relationship as

$$y_n = K_{n,L} \cdot X_L \quad (1.1)$$

- We can begin by weighting the equation, if desired

$$W_{n,n} \cdot y_n = W_{n,n} \cdot K_{n,L} \cdot X_L \quad (1.2)$$

- Then multiplying by,  $K^T$  we obtain

$$K_{L,n}^T \cdot W_{n,n} \cdot y_n = K_{L,n}^T \cdot W_{n,n} \cdot K_{n,L} \cdot X_L \quad (1.3)$$

- And the solution is

$$X_L = [K_{L,n}^T \cdot W_{n,n} \cdot K_{n,L}]^{-1} \cdot K_{L,n}^T \cdot W_{n,n} \cdot y_n \quad (1.4)$$

- Note that,  $X_L = K_{L,n}^{-1} \cdot y_n$  is the direct solution but for a non-square matrix  $K^{-1} \equiv [K^T \cdot K]^{-1} \cdot K^T$ .

## Summary of Geophysical Products, $X_L$

$T(p)$	vertical temperature profile
$q(p)$	vertical water vapor profile ( $\approx 8$ g/kg @ surface)
$L(p)$	vertical liquid water profile (f/ AMSU/HSB)
$O_3(p)$	vertical ozone profile ( $\approx 8$ ppmv @ 6 mb)
$T_s$	surface temperature
$\epsilon(\nu)$	spectral surface emissivity, ( <i>e.g.</i> , 0.95 @ 800 $\text{cm}^{-1}$ )
$\rho_{\odot}(\nu)$	spectral surface reflectivity of solar radiation
$P_{\text{cld}}$	cloud top pressure for $\leq 2$ cloud levels
$\alpha_{\text{cld,fov}}$	cloud fraction for $\leq 2$ cloud levels and 9 FOV's
$CO_2$	total column carbon dioxide ( $\approx 370$ ppmv)
$CH_4(p)$	methane profile ( $\approx 1.65$ ppmv)
$CO(p)$	carbon monoxide profile ( $\approx 0.11$ ppmv)

### Ancillary Information Needed for Retrieval

$P_s$	surface pressure (f/ AVN forecast)
$\theta$	satellite zenith angle
$\theta_{\odot}$	solar zenith angle
$\epsilon_{\text{cld},\nu} \equiv 1$	spectral cloud emissivity for $\leq 2$ cloud levels

## The Cost Function

The idea of minimization of a cost function for the solution of linear equations dates back to Gauss (c. 1800). Most methods, include the AIRS Science Team method, minimize a cost function of the form:

$$\begin{aligned}
 J = & (f_n(X_L) - y_n)^T \cdot N_{n,n}^{-1} \cdot (f_n(X_L) - y_n) \\
 & + (X_L - X_L^0)^T \cdot H_{L,L} \cdot (X_L - X_L^0)
 \end{aligned} \tag{1.5}$$

We find the solution of  $\frac{\partial J}{\partial X_L} = 0$ . Since derivatives of the forward model are a function of the parameters, this problem is inherently non-linear. Therefore, we must iterate. For iteration= $i$  the solution is given by

$$K_{n,L}^i \equiv \left. \frac{\partial f_n(X_L)}{\partial X_L} \right|_{X_L^i} \tag{1.6}$$

$$\begin{aligned}
 X_L^{i+1} = & X_L^0 + \left[ K_{L,n}^{T^i} \cdot N_{n,n}^{-1} \cdot K_{n,L}^i + H_{L,L} \right]^{-1} \cdot K_{L,n}^{T^i} \cdot N_{n,n}^{-1} \cdot \\
 & \left[ y_n - f_n(X_L^i) + K_{n,L}^i (X_L^i - X_L^0) \right]
 \end{aligned} \tag{1.7}$$

However, this form can also be derived using linear algebra (similar to Eqn. 1.4) or maximum likelihood (Bayesian probability density functions).



## Reminder, What is the Forward Model

- The non-scattering infrared forward model has components from the surface (mostly linear), the reflected solar component (mostly linear), a complicated, but very small, down-welling term, and an atmospheric term of the form of a Fredholm integral equation of the 1<sup>st</sup> kind:

$$f_n(X_L^i) \simeq \int_{\nu} d\nu \cdot \Phi(\nu, \nu_0(n)) \cdot \int_p dp \cdot B_{\nu}(T^i(p)) \cdot \frac{\partial \exp \left( - \int_{z'=\infty}^{z(p)} \sum_i \kappa_i(\nu, X_L^i, p, \dots) dz' \right)}{\partial p} \quad (1.8)$$

- The optical depth for a given gas (denoted by a subscript  $i$ ),  $\kappa_i$ , is a complex interacting function of all the parameters,  $X_L^i$ , such as temperature  $T^i(p)$ , moisture, ozone, etc.
- $\Phi(\nu, \nu_0(n))$  is our instrument response function
- This equation can be highly non-linear for composition (*i.e.*, moisture, ozone, etc.) retrievals.

## Damping and the “Background Term”

- The matrix  $H_{L,L}$  is a form of regularization
  - Prevents inverse from being singular, *i.e.*, it stabilizes the inverse.
  - $X^{i+1} - X^i$  is smaller (*i.e.*, damped) than the LSQ solution (see Eqn. 1.4), which I will denote as  $X^{i+1}(H = 0)$  in these notes.
- The radiances are modified by the right hand term,  $K_{n,L}^i (X_L^i - X_L^0)$ 
  - When  $H$  is non-zero, the part that wasn't believed on the 1<sup>st</sup> iteration,  $X^1(H = 0) - X^1$ , must be subtracted from the radiances.
  - If this term is neglected then the part that wasn't believed would be introduced into subsequent iterations and the effect would be an un-damped LSQ retrieval,  $X^{i+1}(H = 0)$ .
  - This is why, a physical retrieval cannot use the results from a previous physical retrieval  $\Rightarrow$  RULE #1 of Iteration

## Heritage of Regularization Approaches

year	authors	method of regularization
1943	Levenberg	Steepest Descent & Newtonian Iteration
1963	Marquart	Hessian operator, $\mathbf{H} = \Delta \mathbf{y} \cdot \frac{\partial^2 (y_n - f_n(X_L))^2}{\partial X_L^2}$
1963	Twomey, Tikhonov	$\mathbf{H} = \lambda \cdot \mathbf{I}$ , prevent singularities
≈ 1970	Twomey	$\mathbf{H}$ minimizes vertical derivatives, <i>e.g.</i> , $\partial T / \partial z$
1970	Backus, Gilbert	Compute Optimal Vertical Functions f/ sounding
1972	Conrath	Trade-off: Vertical Resolution versus $\chi^2$ error
≈ 1970	Wark & Fleming	Use $\mathbf{H}$ = co-variance as a constraint
1976	Rodgers	Use <i>a posteriori</i> PDF's as a constraint, $\mathbf{H} = \mathbf{S}_a^{-1}$
1989	Eyre	Formalization of forward model errors in $\mathbf{N}_{n,n}^{-1}$
1992	Hanel, Conrath	Optimal functions/vertical resolution by SVD
1992	Hansen	L-curve, finding optimal $\lambda$ via SVD
1996	Phalippou	For $\mathbf{q}(\mathbf{p})$ , use relative humidity as a constraint
1997	Schimpf & Schreier	Use of SVD to determine $\mathbf{H}$
1999	Li	use residuals to derive $\lambda$
2000	Peckham & Grippa	Use lapse rate as a constraint

## Constraints on AIRS Science Team Algorithm

- Retrieval methodology must be able to handle cloud cleared radiances (CCR's)
  - Random noise amplification,  $\frac{1}{3} \leq A < \approx 3$
  - Large spectrally correlated component
  - Statistical *a-priori* difficult to implement
- Retrieval should have minimal sensitivity to first guess.
  - Maximize contribution from instrument radiances.
  - Maximize sensitivity to and understanding of climate signals
  - Trade-off: model background states  $\Rightarrow$  to use or not to use.
- Retrieval should not artificially constrain problem.
  - Minimize sensitivity to incorrect statistics, *e.g.*, in frontal zones avoid statistical damping.
  - Trade-off: Stability versus Impact

## Cloud Clearing Methodology (Chapter 7 of rs\_notes.pdf)

This cloud clearing methodology has a long heritage starting from the original papers (Smith, 1968, Chahine, 1974), Chahine, 1975, Chahine, 1977, Chahine *et al.* 1977, McMillin and Dean 1982, Smith *et al.* 1992, Susskind, *et al.* 1998, Joiner and Rokker 2002, Susskind, *et al.* 2003. The fundamental features of the AIRS cloud clearing algorithm are

- Use the  $J = 9$  AIRS cloud scenes,  $\mathbf{R}_{n,j}$ , without any *a-priori* constraint, such as preferential grouping, to compute the extrapolation parameters,  $\boldsymbol{\eta}_j$ .

$$\hat{\mathbf{R}}_n = \langle \mathbf{R}_{n,j} \rangle_j + (\langle \mathbf{R}_{n,j} \rangle_j - \mathbf{R}_{n,j}) \cdot \boldsymbol{\eta}_j \quad (1.9)$$

- Determine the number of cloud *formations* and constrain the number of degrees of freedom for solution of  $\boldsymbol{\eta}_j$  to the number of cloud formations.
- Compute both CCR's and error estimates for the CCR's,  $\boldsymbol{\delta}\hat{\mathbf{R}}_n\boldsymbol{\delta}\hat{\mathbf{R}}_n$ , specifically taking into account the noise amplification induced by the linear extrapolation and the spectrally correlated component of the radiance error due to error covariance of the  $\boldsymbol{\eta}$ 's,  $\boldsymbol{\delta}\boldsymbol{\eta}_j\boldsymbol{\delta}\boldsymbol{\eta}_j^T$ .
- Compare the clear state estimate with the AIRS retrieval products and reject cases that violate any of the assumptions of cloud clearing.

## A non-Traditional Look at the Cost Function, 1/3

$$\begin{aligned} \mathbf{J} = & \left( \mathbf{f}_n(\mathbf{X}_L^i) - \mathbf{y}_n \right)^T \cdot \mathbf{N}_{n,n}^{-1} \cdot \left( \mathbf{f}_n(\mathbf{X}_L^i) - \mathbf{y}_n \right) \\ & + \left( \mathbf{X}_L^i - \mathbf{X}_L^0 \right)^T \cdot \mathbf{H}_{L,L} \cdot \left( \mathbf{X}_L^i - \mathbf{X}_L^0 \right) \end{aligned} \quad (1.10)$$

- We can compute the error in cloud cleared observations,  $\delta \hat{\mathbf{R}}_n \delta \hat{\mathbf{R}}_n$  very well.
- We can estimate errors in the forward model. For parameters held constant,  $\mathbf{X}_b$ , the obs-calc error covariance is

$$\mathbf{N}_{n,n'}^{-1} = \mathbf{K}_{n,b} \cdot \delta \mathbf{X}_b \delta \mathbf{X}_{b'}^T \cdot \mathbf{K}_{b'n'}^T + \delta \hat{\mathbf{R}}_n \delta \hat{\mathbf{R}}_n \quad (1.11)$$

- *a-priori* information enters the system through statistical estimates of  $\delta \mathbf{X}_b \delta \mathbf{X}_{b'}^T$ 
  - In the sense of estimates of errors in  $\mathbf{X}_L^0$
  - In the sense of null space errors, the minimum allowed value of  $\delta \mathbf{X}_b \delta \mathbf{X}_{b'}^T$
- But, we also compute the formal errors of the solution,  $(\delta \mathbf{X}_L \delta \mathbf{X}_L^T)^s$  for each retrieval step =  $s$  (details in Section 21.9 of rs\_notes.pdf).

## A non-Traditional Look at the Cost Function, 2/3

- $(\delta X_L \delta X_L^T)^s$  from step =  $s$  becomes the  $(\delta X_b \delta X_b^T)^{s+1}$  in step =  $s + 1$ , *e.g.*, we solve for  $T(p)$  and  $\delta T(p) \delta T(p)^T$  and then use that error covariance when we solve for  $q(p)$ ,  $O_3(p)$ , etc., in later steps.
- Therefore, an improvement in temperature errors, for example, can be used to improve our moisture retrieval (vice a versa).
- This formulation also brings spectral correlation (*i.e.*, *a priori* knowledge via the forward calculation) into the solution via  $K_{n,b}(X_L^i)$  on a case-by-case basis.
  - Spectral correlation is a function of other state parameters. For example, temperature lapse rate changes sensitivity of all the composition derivatives.
  - $N$  equations of  $y_n$  changed into  $N$  new equations:  $N_{n,n'} \cdot y_{n'}$ .
  - This is a powerful concept that when used properly
    - \* allows separation of mixed signals, *e.g.*,  $T(p)$  and  $\text{CO}_2$
    - \* minimizes sensitivity to biases, *e.g.*, surface effects ( $T_{skin}$ ,  $P_{surf}$ , emissivity) in  $T(p)$  retrieval.

## A non-Traditional Look at the Cost Function, 3/3

- We can compute  $H$  from information content of  $(K^T N^{-1} K)$  by singular value decomposition (SVD, a.k.a. empirical orthogonal functions (EOF's) see Section 21.3, rs\_notes.pdf)

$$\Lambda_{k,k}^i \equiv U_{k,L}^{T^i} \left[ K_{L,n}^{T^i} \cdot (N_{n,n}^s)^{-1} \cdot K_{n,L}^i \right] \cdot U_{L,k}^i. \quad (1.12)$$

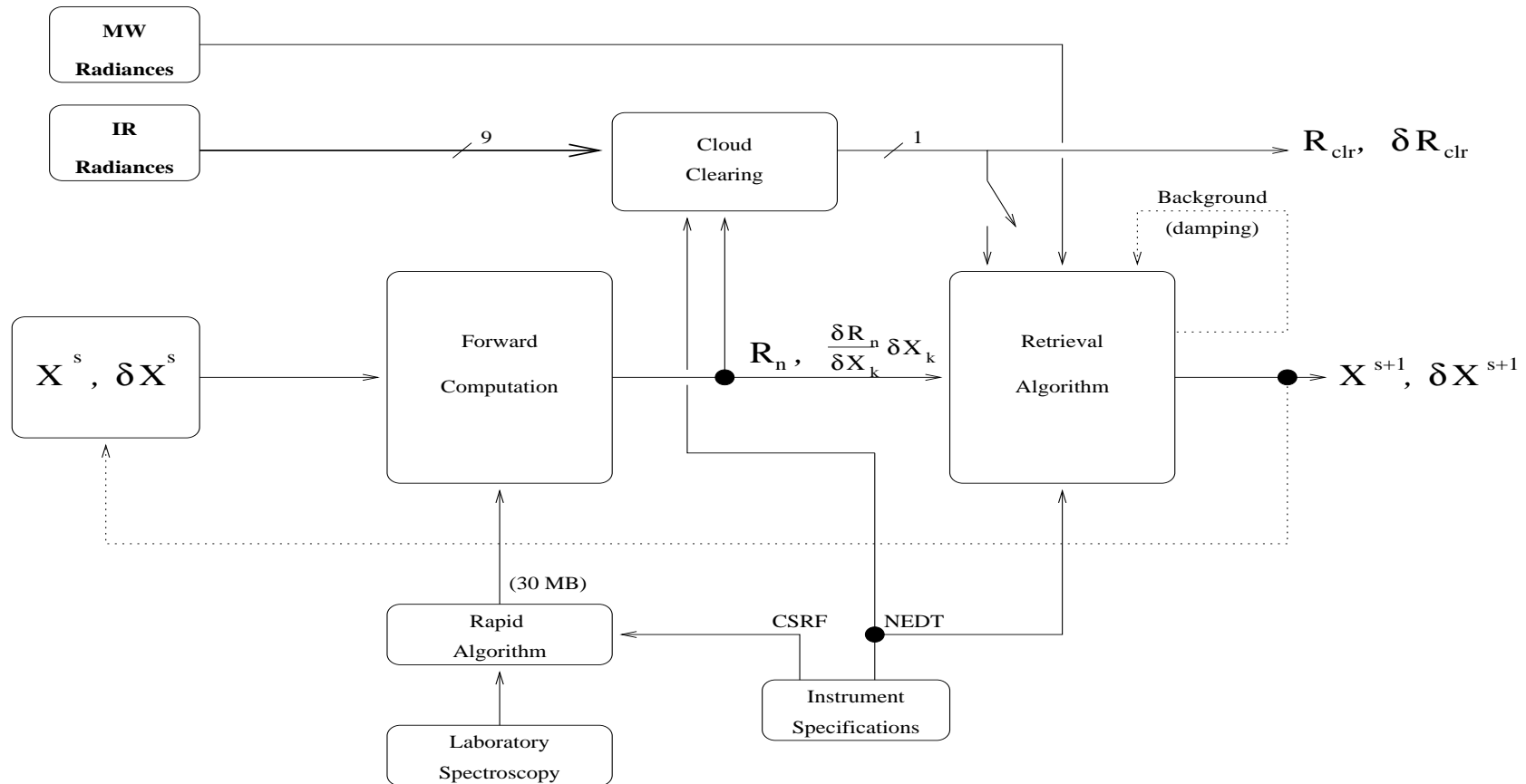
- $\Lambda_{k,k}^i$  is diagonal with elements equal to  $\lambda_k$
- When  $\lambda_k^i \gg 1$  the terms are well determined.  $K_{n,L}^i \cdot U_{L,k}^i$  are new Jacobians with very high signal to noise:  $H = 0$
- When  $\lambda_k^i \rightarrow 0$  the observations have no influence on the solution:  $H \rightarrow \infty$  and components of  $X_L^{i+1} \rightarrow X_L^0$
- When  $\lambda_k^i$  is small and significant we add a  $\Delta\lambda_k^i$  (details in Section 21.4 of rs\_notes.pdf) which is equivalent (see Section 21.8, rs\_notes.pdf) to adding a case dependent  $H^i$  given by

$$H_{L,L'}^{s,i} = U_{L,k}^i \cdot \Delta\Lambda_{k,k}^i \cdot U_{k,L}^{T^i} \quad (1.13)$$

- SVD determines the optimal fraction of the *a priori* information to use..
- We think this is more robust than using ensemble statistics of  $I$  cases to compute a static *a priori* covariance  $S_a \equiv \delta X_{L,i} \delta X_{i,L}^T$



# Overview of the AIRS Multi-spectral Physical Retrieval System



The atmospheric state,  $X_L^s$ , and the error estimate of that state,  $\delta X_L^s$ , are used to minimize the residuals in observed minus computed radiances in each retrieval step= $s$ .



It feels like this, eh?

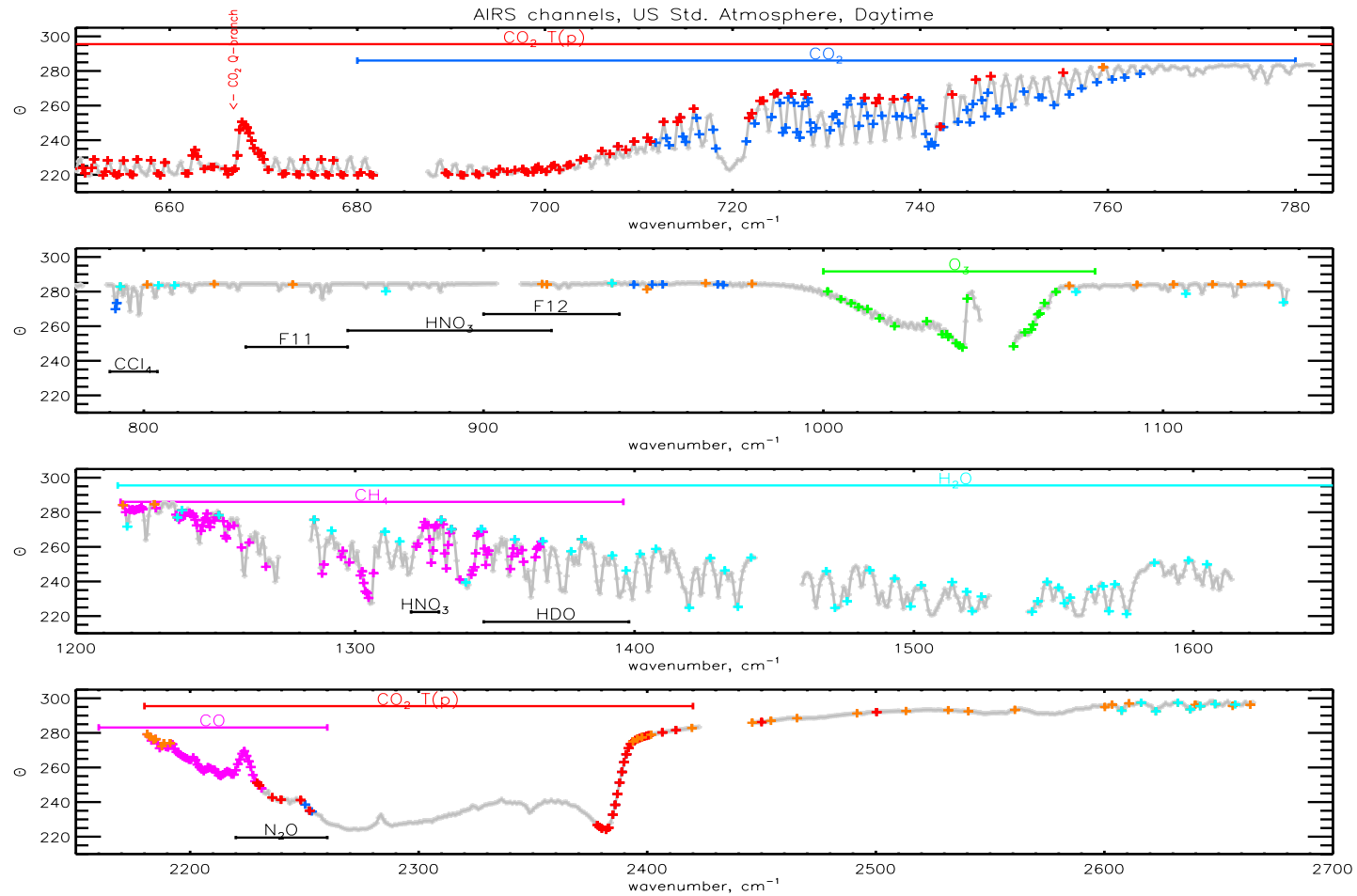
## Philosophy of AIRS *Team* Algorithm

- Utilize microwave-only product (MIT maximum likelihood algorithm  $\Rightarrow T(p), q(p), T_{skin}$ ) to estimate the infrared clear radiance estimate for initial cloud clearing.
- Utilize eigenvector regression (NOAA/NESDIS) to provide the first guess state for the physical algorithm. This solution contains the fine vertical structure information based on  $\approx 1600$  AIRS channels..
- Utilize a physical retrieval (NOAA/GSFC) to improve the state.
  1. Microwave and infrared observations are used in each step.
  2. Use microwave observations and products to reject cases with poor cloud clearing.
    - Reject if Obs-Calc of coupled retrieval is too large
    - Perform microwave-only retrieval using coupled retrieval as first guess. Reject if  $\Delta T(p)$  is too large.
  3. All observations are used at their native angles of observation and the forward model is computed at the correct angles.
  4. For cloud clearing we do perform a local angle correction of  $1.1^\circ$  with each  $3 \times 3$  set of FOV's

## Philosophy of AIRS Physical Algorithm

1. Embed an information content analysis into each step to determine the optimal damping (regularization) for each case.
  - Cloud cleared radiances are both case and iteration dependent.
  - Propagate a formal geophysical error estimate through each step.
  - Compute an estimate of the *a-priori* covariance at each step.
2. Take advantage of parameters that are separable (*i.e.*, pay attention to spectroscopy and radiative transfer features in our spectrum)
  - For example, we can solve for  $T_{surf}$  holding all other variables (*e.g.*, water) constant, since  $T_{surf}$  is quite linear. Steps are denoted as a superscript  $s$  in my notes.
  - BUT, if a step is repeated (*e.g.*, when an error estimate has been improved) NEVER use the products from the previous step.
3. Select channels that are “spectrally pure”, that is
  - Have a high sensitivity to what is being solved for.
  - Have a low sensitivity to the parameters held constant (*i.e.*, keep our error estimates small when separating variables)
4. Use optimal number of parameters for each retrieval step are determined in simulation (Backus-Gilbert optimization) to speed up processing.

# Channels used in the AIRS retrieval algorithm



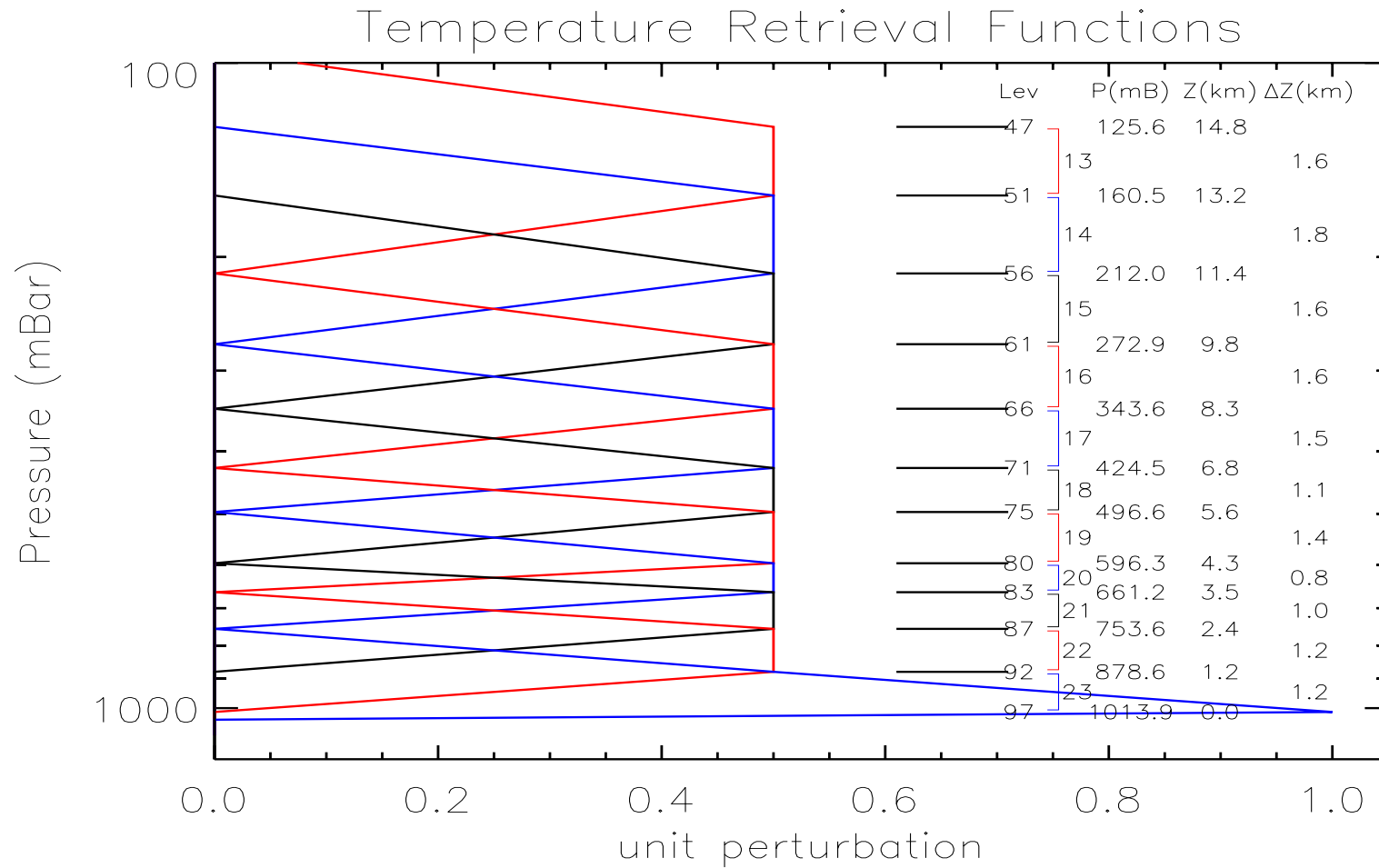
## Justification for Vertical and Spectral Functions

- A fine vertical grid is required for accurate computation of the absorption coefficients,  $\kappa_i(\nu, p(z), X^{s,i}, \theta)$  and radiances.
- There are about 85 independent pieces of information in AIRS (2378 chl's), IASI (8461 chl's), and CrIS (1305 chl's)
- Solving for 100 vertical levels of T(p), q(p), O<sub>3</sub>(p), CO(p), CH<sub>4</sub>(p), and CO<sub>2</sub>(p) wastes time and can destabilize the retrieval.
- If we allowed 2378 emissivities, there would be nothing left to solve for.
- Functions,  $F_{L,j}^s$ , and associated parameters,  $A_j^{s,i}$ , are chosen in a trade-off between resolution and stability for each retrieval step. Analogous to Backus & Gilbert (1970) trade-off. Also discussed in Hanel, (1992).
- Major issue for code execution time and improves stability.

## Formulation of Vertical and Spectral Functions

- Temperature functions are additive vertical trapezoids.
- Composition functions are multiplicative vertical trapezoids.
  - Radiance kernel is  $\propto \exp(\kappa(X_L^{s,i}))$ ,
  - $\kappa(X_L^{s,i})$ , is the optical depth  $\propto X_L^{s,i}$ .
  - Therefore, composition variables are more linear in  $\ln(X_L^{s,i})$
  - $\partial \ln(X_L^{s,i}) \propto \frac{\partial X_L^{s,i}}{X_L^{s,i}}$  which is a % change in  $X_L^{s,i}$ .
- Emissivity functions are additive spectral triangles.
- A scaling parameter  $\hat{A}_j^s$  is used to create dimensionless parameters and adjust scale between different functional groups (*e.g.*, when mixing T(p), q(p), and emissivity in one retrieval).
- The Jacobian,  $K_{n,L}^{s,i}$ , becomes a set of new derivatives,  $S_{n,j}^{s,i}$ , in which groups of parameters in  $L$  space are grouped together in  $J$  space.
- Sub-sets (*e.g.*, temperature) of vertical and spectral functions must sum to unity:  $\sum_j (F_{L,j}^s) = 1$  for a group of functions.

# Example of functions, $F_{L,j}$ , for $T(p(L))$ retrieval





## Kernel Function replaced by Sensitivity Matrix

- For additive functions the S-matrix is given by

$$S_{n,j}^{s,i} \equiv \frac{\partial f_n \left( X_L^{s,i} + F_{L,j}^s \cdot A_j^s \right)}{\partial A_j^s} \cdot \Delta \hat{A}_j^s \quad (2.1)$$

$$\simeq f_n \left( X_L^{s,i} + F_{L,j}^s \cdot \hat{A}_j^s \right) - f_n \left( X_L^{s,i} \right) \quad (2.2)$$

- For multiplicative functions the S-matrix is given by

$$S_{n,j}^{s,i} \equiv \frac{\partial f_n \left( X_L^{s,i} \cdot \left( 1 + F_{L,j}^s \cdot A_j^s \right) \right)}{\partial A_j^s} \cdot \Delta \hat{A}_j^s \quad (2.3)$$

$$\simeq f_n \left( X_L^{s,i} \cdot \left( 1 + F_{L,j}^s \cdot \hat{A}_j^s \right) \right) - f_n \left( X_L^{s,i} \right) \quad (2.4)$$

- Analytic derivatives on the RT grid do not help our algorithm,  $\delta$  function perturbations are sub-optimal (Backus+Gilbert).
- Single sided finite difference is currently used, we will explore the benefit of double-sided and dynamically scaled derivatives someday. This is not our biggest error source!!!

So ... we really solve for dimensionless parameters

The minimization equation for both additive and multiplicative forms is given by

$$\begin{aligned}
\Delta A_j^{s,i+1} &\equiv A_j^{s,i+1} - A_j^{s,0} \\
&= \left[ \left( S_{j,n}^{T^{s,i}} \right) \cdot \left( N_{n,n}^s \right)^{-1} \cdot S_{n,j}^{s,i} + H_{j,j}^{s,i} \right]^{-1} \cdot S_{j,n}^{T^{s,i}} \cdot \left( N_{n,n}^s \right)^{-1} \cdot \\
&\quad \left[ y_n - f_n(X_L^{s,i}) + S_{n,j}^{s,i} \left( A_j^{s,i} - A_j^{s,0} \right) \right]
\end{aligned} \tag{2.5}$$

And to return to parameter  $L$  space is done by combining the components and dividing by the scaling factors,  $\hat{A}_j^s$ .

$$\Delta X_L^{s,i+1} = \sum_j F_{L,j}^s \cdot \left( \Delta A_j^{s,i+1} \cdot I_{j,j} \cdot \Delta \hat{A}_j^{-1^s} \right) \quad \text{Additive} \tag{2.6}$$

$$\Delta X_L^{s,i+1} = \prod_j \left( 1 + F_{L,j}^s \right) \cdot \left( \Delta A_j^{s,i+1} \cdot I_{j,j} \cdot \Delta \hat{A}_j^{-1^s} \right) \quad \text{Multiplicative} \tag{2.7}$$

## Finally, we regularize

- We perform the information content analysis on our new functions

$$\Lambda_{k,k}^{s,i} \equiv U_{k,j}^{T^{s,i}} \cdot \left[ S_{j,n}^{T^{s,i}} \cdot \left( N_{n,n}^s \right)^{-1} \cdot S_{n,j}^{s,i} \right] \cdot U_{j,k}^{s,i} \quad (2.8)$$

- This transformation is equivalent to new Jacobians,  $S_{n,j}^{s,i} \cdot U_{j,k}^{s,i}$
- We determine  $\Delta\lambda$  as follows (see Section 21.8, rs\_notes.pdf)

$$\Delta\lambda_k^{s,i} = 0 \quad \text{for } \lambda_k^{s,i} \geq \lambda_c^s \quad (2.9)$$

$$= \sqrt{\lambda_k^{s,i}} \cdot \left( \sqrt{\lambda_c^s} - \lambda_k^{s,i} \right) \quad \text{for } \lambda_k^{s,i} < \lambda_c^s \quad (2.10)$$

$$= \infty \quad \text{for } \lambda_k^{s,i} \leq (0.05)^2 \cdot \lambda_c^s \quad (2.11)$$

## And Solve for Our Parameters

- We solve for the new parameters

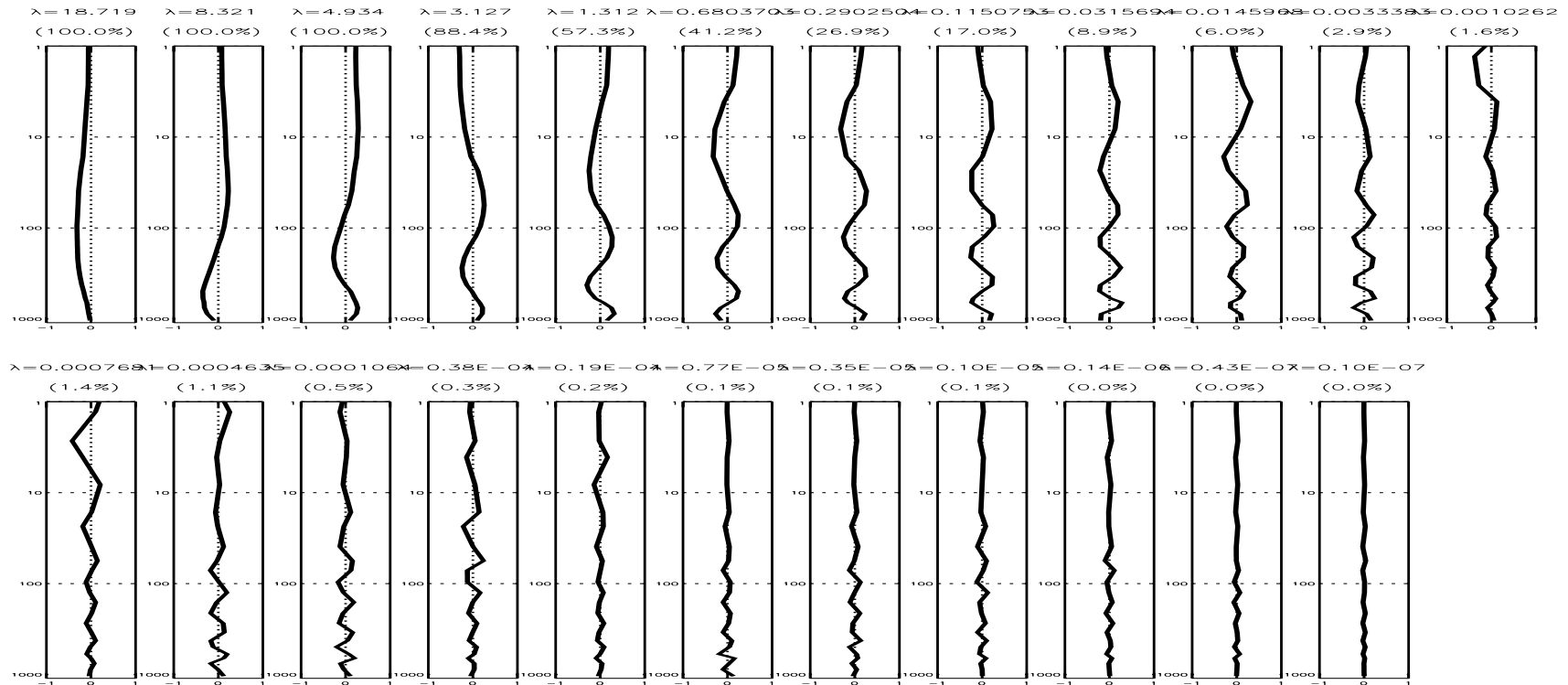
$$\Delta A_j^{s,i+1} = U_{j,k}^{s,i} \cdot \Delta B_k^{s,i+1} \quad (2.12)$$

$$\Delta B_k^{s,i+1} = \left( \frac{1}{\lambda_k^{s,i} + \Delta \lambda_k^{s,i}} \right) \cdot U_{k,j}^{Ts,i} \cdot S_{j,n}^{Ts,i} \cdot (N_{n,n}^s)^{-1} \cdot \left[ y_n - f_n(X_L^{s,i}) + S_{n,j}^{s,i} (A_j^{s,i} - A_j^{s,0}) \right] \quad (2.13)$$

- Note that the change in the parameters,  $\Delta A_j^{s,i+1}$  associated with original functions,  $F_{L,j}^s$  is equivalent to a transformed parameter change,  $\Delta B_k^{s,i+1}$ , associated with a new function  $G_{L,k}^{s,i} = F_{L,j}^s \cdot U_{j,k}^{s,i}$ .
- It is illustrative to visualize the transformed functions.

## Example of SVD *optimal* $T(p)$ Functions

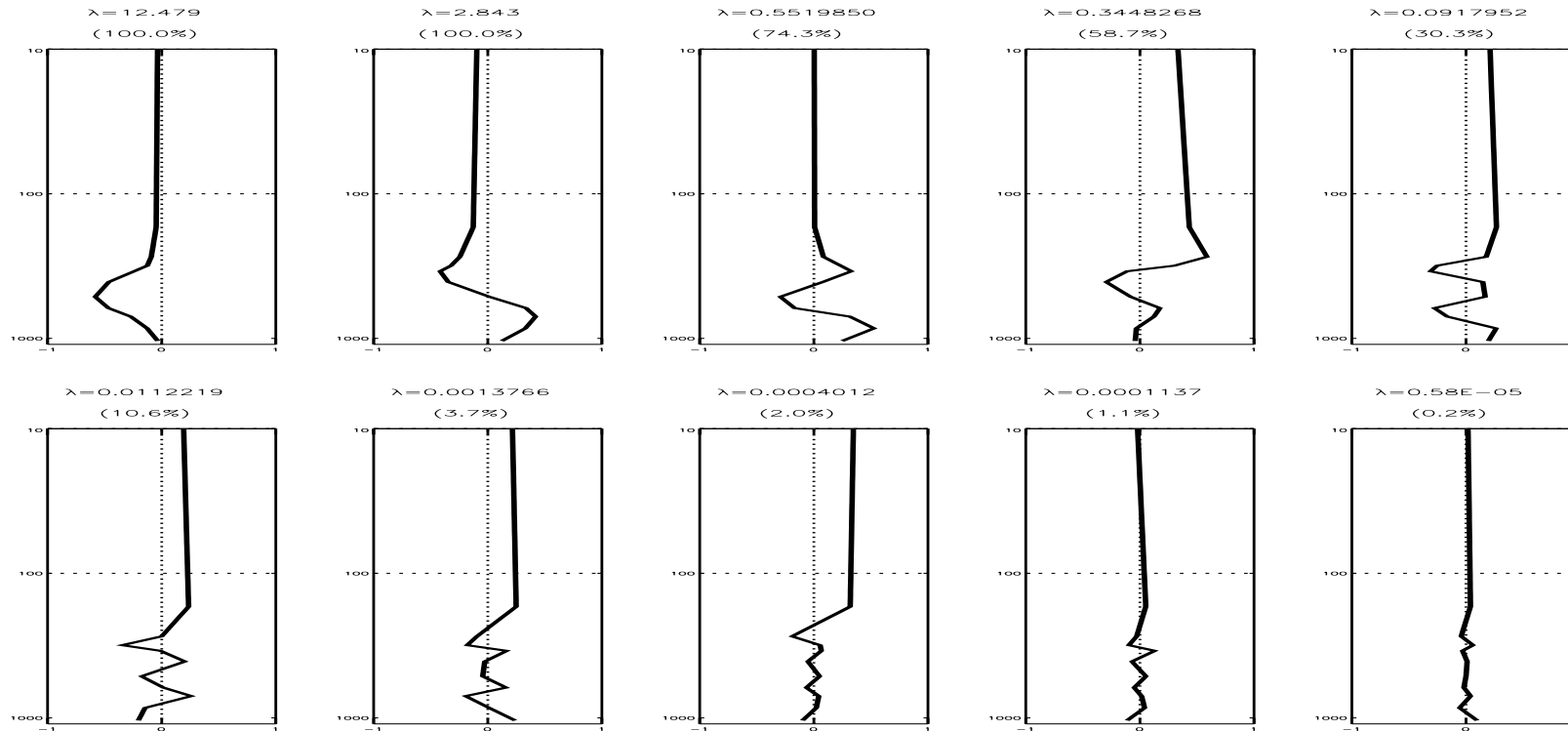
: Profile 977 Temperature #1/2 Eigenfunctions,  $\lambda_m = 4.0000$



For temperature functions, the new set of vertical functions,  $F_{L,j}^s \cdot U_{j,k}^{s,i}$  are shown for the AIRS temperature retrieval information content analysis. In this case, ten functions of the 23 functions are determined to better than 5%.

## Example of SVD *optimal* $q(p)$ Functions

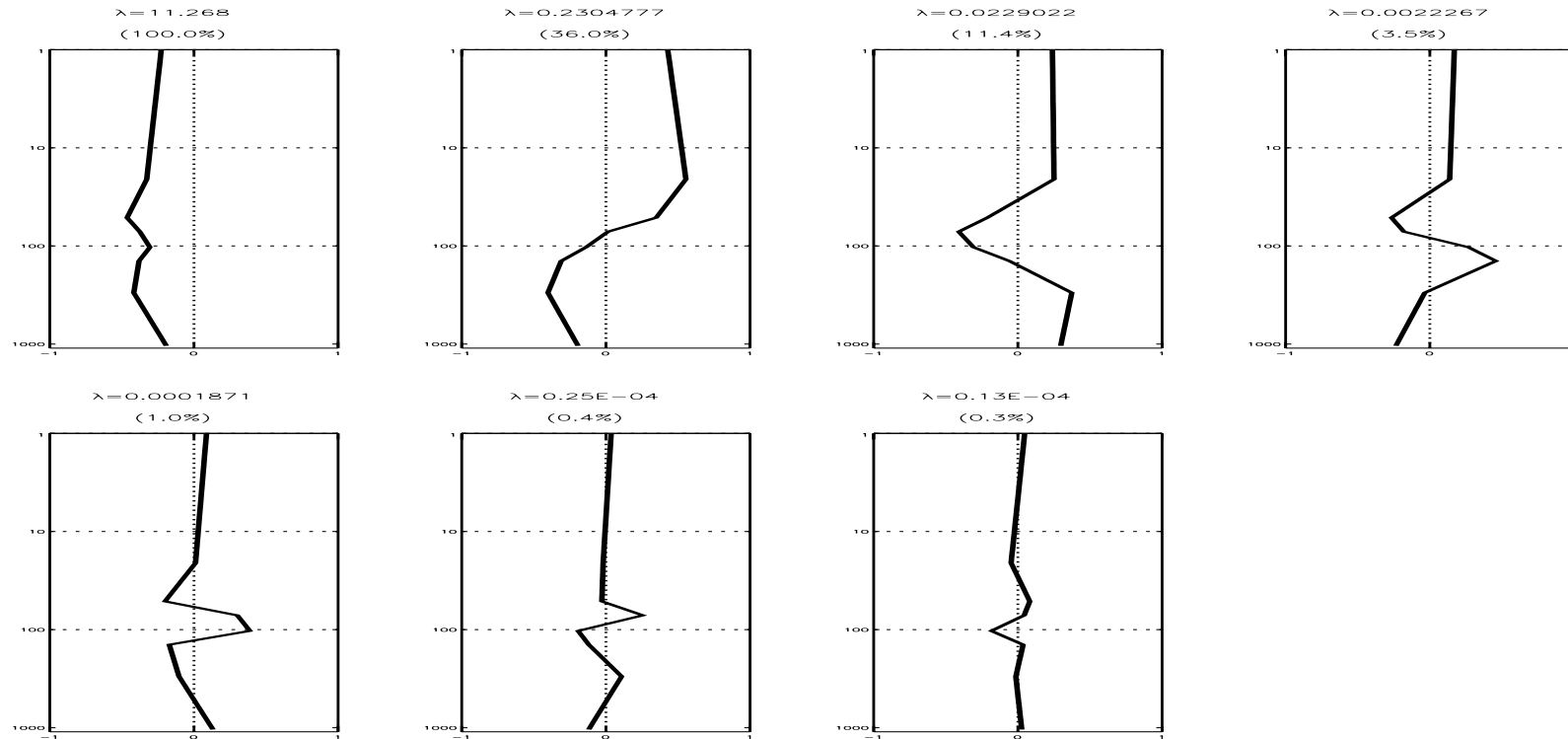
: Profile 977 Water Vapor (IR/MHS) Eigenfunctions,  $\lambda_m = 1.0000$



*For water functions, the new set of vertical functions,  $F_{L,j}^s \cdot U_{j,k}^{s,i}$  are shown for the AIRS temperature retrieval information content analysis. In this case, six functions of the ten functions are determined to better than 5%.*

## Example of SVD *optimal* $O_3(p)$ Functions

: Profile 977 Ozone Eigenfunctions,  $\lambda_m = 1.7778$



For ozone functions, the new set of vertical functions,  $\mathbf{F}_{L,j}^s \cdot \mathbf{U}_{j,k}^{s,i}$  are shown for the AIRS temperature retrieval information content analysis. In this case, three functions of the seven functions are determined to better than 5%.

## Propagation of Formal Errors

- The uncertainty in  $\Delta B_k^{s,i+1}$  ( $\Delta \lambda_k^{s,i} = 0$ ) are uncorrelated and equal to  $(\lambda_k^{s,i})^{-\frac{1}{2}}$ .
- The fraction of  $\Delta B_k^{s,i}$  solved for is equal to  $\phi_k^{s,i} = \lambda_k^{s,i} / (\lambda_k^{s,i} + \Delta \lambda_k^{s,i})$
- The propagated error in  $B$  space is the RSS of the transformed first guess error,  $\delta B^{s,0}$ , and the error from the radiances.

$$\delta B_k^{s,i+1} = \sqrt{\left( (1 - \phi_k^{s,i}) \cdot \delta B_k^{0,s} \right)^2 + \left( \phi_k^{s,i} \cdot \frac{1}{\sqrt{\lambda_k^{s,i}}} \right)^2} \quad (2.14)$$

- In  $\Delta A$  space the errors are correlated and can be computed from

$$(\delta A_j \delta A_j)^{s,i+1} = U_{j,k}^{s,i} \cdot (\delta B_k^{s,i})^2 U_{k,j}^{T,s,i} \quad (2.15)$$

$$\delta A^{s,i+1}(j) \simeq \sqrt{\sum_k U^{s,i}(j, k)^2 \cdot (\delta B^{s,i+1}(k))^2} \quad (2.16)$$



## Propagation of Formal Errors

- The errors in the geophysical products are computed in the root-sum-squared (RSS) sense from parameter errors:

$$(\delta X_L \delta X_L^T)^{s,i+1} = F_{L,j}^s \cdot (\delta A_j \delta A_j)^{s,i+1} \cdot F_{j,L}^{Ts} \quad (2.17)$$

$$\delta X^{s,i+1}(L) \simeq \sqrt{\sum_j \delta (A^{s,i}(j))^2 \cdot F^s(j, L)^2} \quad (2.18)$$

- These error estimates can be used to compute  $(N_{n,n}^s)^{-1}$  and, therefore, propagated into the next retrieval step; however, in practice we only keep the diagonal components of these errors.
- It is a bit more complicated than this, due to handling of null space errors,  $\Rightarrow$  but I figure we have ALL had enough equations.

## Some Post-Launch Issues



Now we have the theory, we are ready to work with real data.

## Issues with AMSU's Estimate of CLEAR State

- Microwave side-lobe corrections (SLC's) for the Aqua platform are more complex than the POES platforms and have NOT been applied to date.
- A large microwave tuning (empirical side-lobe correction) has been employed to mitigate SLC issues.
- A poor AMSU first guess has a negative impact on cloud clearing and, therefore, all AIRS products.
- To understand the impact to AIRS products, we are using a model analysis to increase our information content.
  1. To assess the impact of AMSU SLC issues on the AIRS products.
  2. To assess the need for tuning and/or RTA improvements.
- We are building the ability to bring in MODIS clear pixels co-located to AIRS FOV's to improve our QA & information content.

## TUNING and MODEL ERROR TERMS

For discussion, assume a retrieval equation looks like

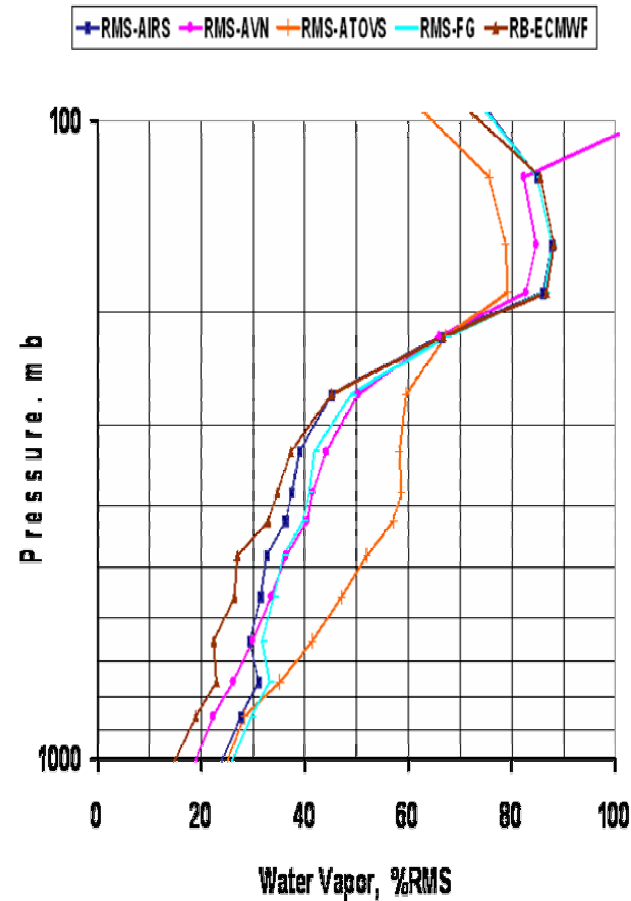
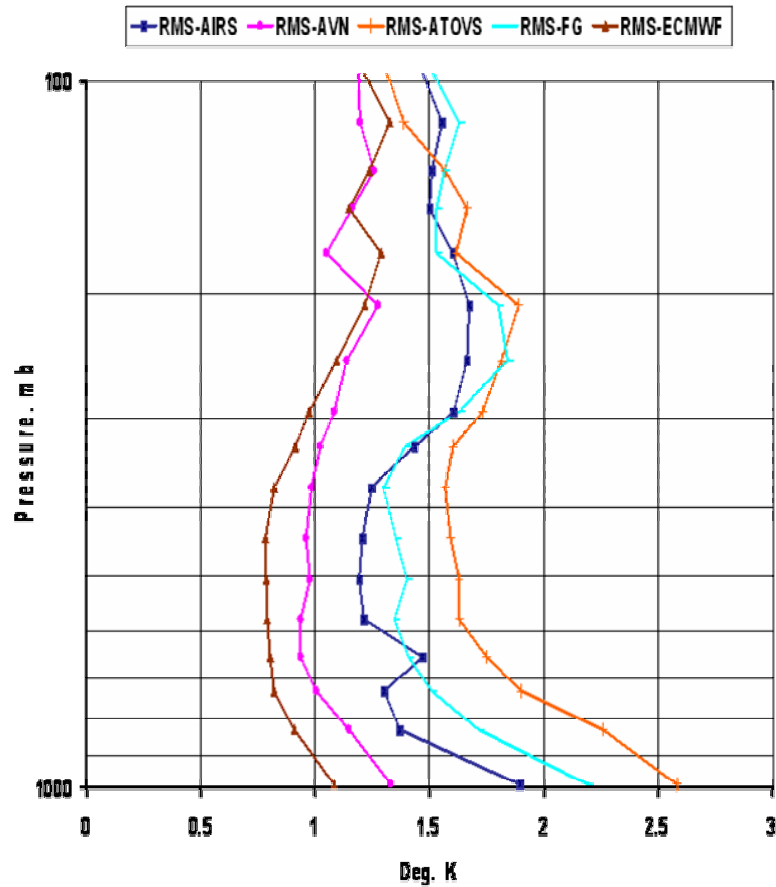
$$\Delta X_j^{s,i+1} = \left[ S_{j,n'}^{T^{s,i}} \cdot \left( N_{n',n}^s \right)^{-1} \cdot S_{n,j}^{s,i} + H_{j,j}^{s,i} \right]^{-1} \cdot S_{j,n'}^{T^{s,i}} \cdot \left( N_{n',n}^s \right) - 1 \cdot \left[ y_n - f_n(X_L^{s,i}) + \Psi_n^{s,i} + T(n) \right] \quad (2.19)$$

- $S_{n,j}^{s,i}$  is the sensitivity of channel  $n$  to parameter  $A_j^s$
- $\Psi_n^{s,i}$  is the background term derived from *a-priori* contribution.
- $T(n)$  is radiance tuning, if applied.

With real data we have other error sources, such as rapid transmittance algorithm (RTA) and spectroscopy errors that we can write as  $E_{n',n}$ .

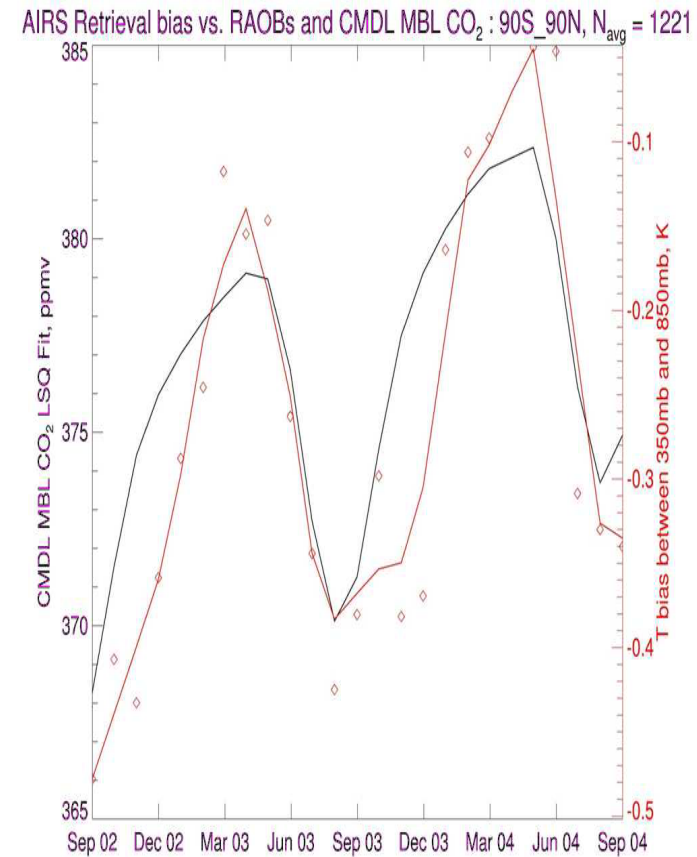
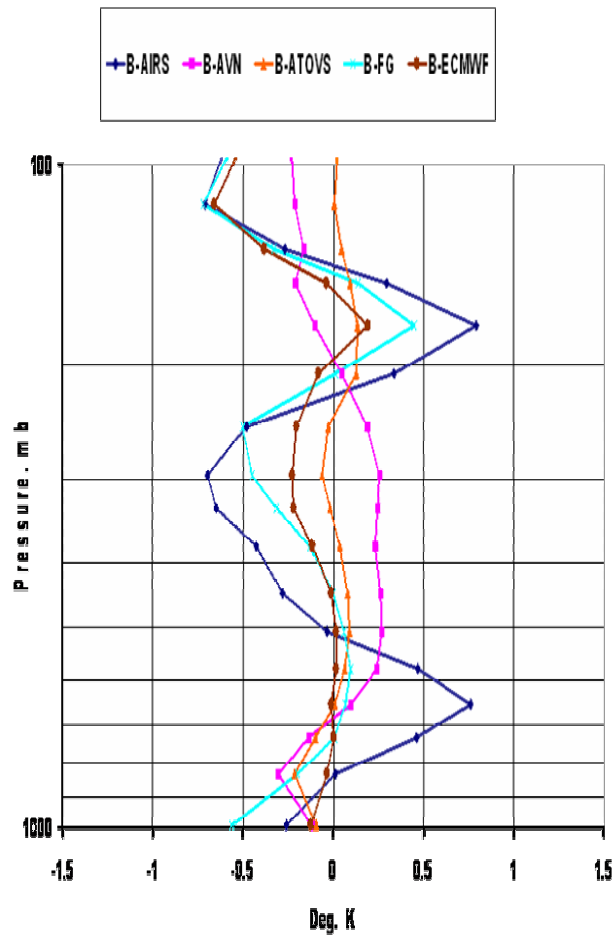
$$N_{n',n}^s = N_{n',n}^s + E_{n',n} \quad (2.20)$$

# Example of RMS Statistics versus RAOB's 32,000 co-located cases from Sep. 2002 to Sep. 2004



These figures provided by Murty Divakarla.

## Example of T(p) BIAS Statistics versus RAOB's 32,000 co-located cases from Sep. 2002 to Sep. 2004

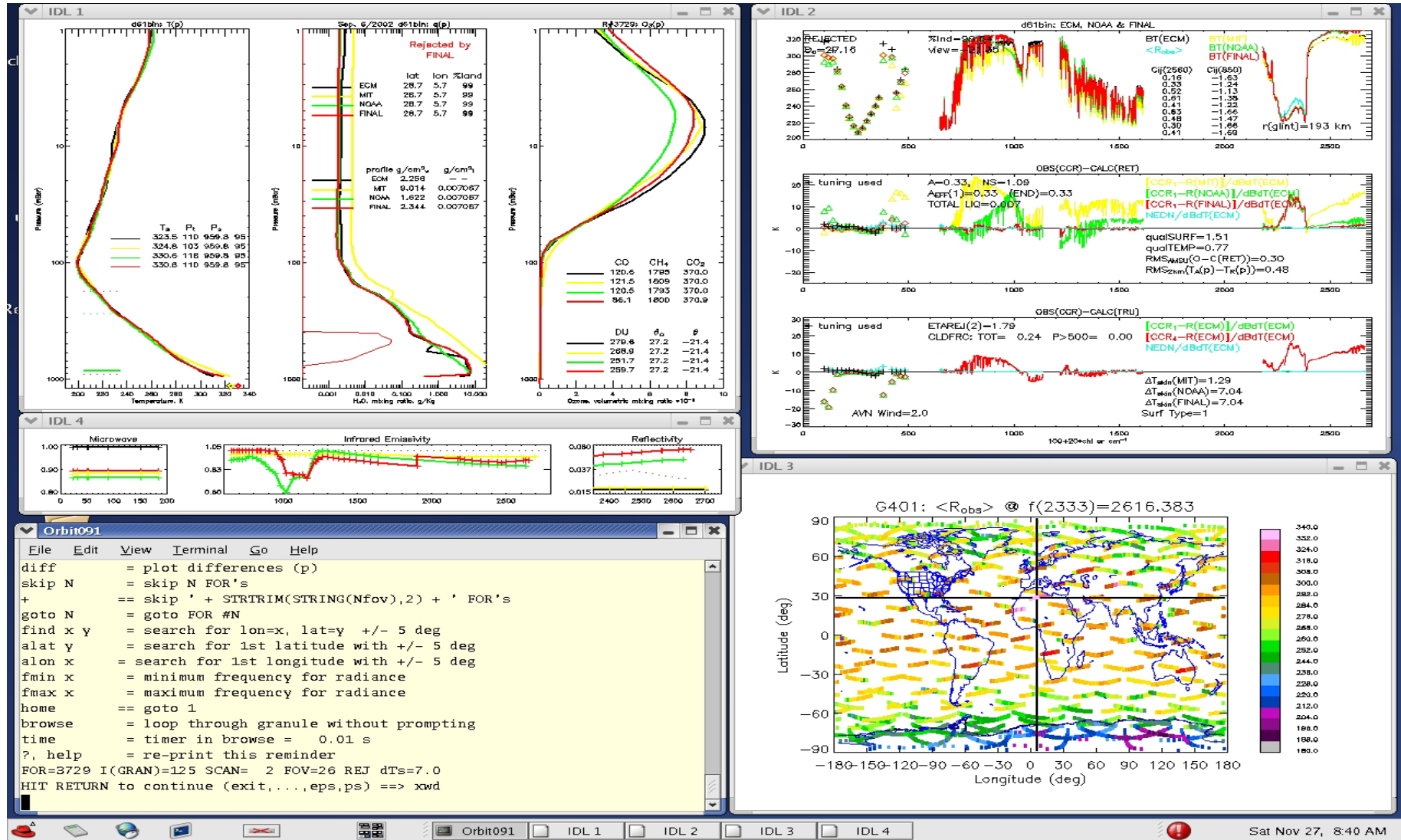


These figures provided by Murty Divakarla & Eric Maddy.

## Code Development

- There is one code for AIRS, IASI, and CrIS.
- The retrieval system code is a total of about 106,000 lines of FORTRAN code.
- We have simulation capability (another 24,000 lines of code).
  - Can perform instrument trade studies.
  - Can study retrieval theory in simulation.
  - Simulation can be built from models and/or from AIRS retrieval products.
- Scientific evaluation is built into the code.
  - Each retrieval case can be compared to a reference state (AVN forecast, ECMWF forecast, RAOB, or truth (in simulation)) at every single iteration and step.
  - Diagnostics exist in radiance space and geophysical space.
- There are over 100,000 lines of IDL code for display and analysis of diagnostic output.

# Example of Diagnostic Capabilities





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